

- I. Axial Tilt and Orbit around Sun—
 - A. Earth tilts 23.5° to orbit around Sun
 - B. Latitudes of Sun incidence
 1. polar circles—Arctic, Antarctic
 2. tropical circles—Cancer, Capricorn
 3. equator
 - C. Days corresponding to Sun incidence
 1. solstice—overhead Sun at tropical circle—June 22, December 22
 2. equinox—overhead Sun at equator—March 22, September 22
 - D. Day length varies with latitude and season, due to circle of illumination
 1. longest/shortest at solstice, at high latitudes
 2. equinox—equal day and night all over the world
 - E. Heating
 1. Greater heating capability with higher Sun angle
 - a. latitude controls on Sun energy
 - b. seasonal controls on Sun energy
 2. Low latitudes have high Sun angle year round
 - F. Incoming Solar Radiation
 1. Temperature Variation greater in areas closer to poles
 2. Temperature Distribution—affected by land-water relation as well as latitude
- II. Weather and climate
 - A. Controlled by two predominant factors
 1. latitude
 2. land-water relations
 - B. creates patterns of
 1. temperature
 2. humidity
 3. cloudiness
 4. precipitation
 5. pressure and wind patterns
- III. Composition of atmosphere
 - A. mixture of gases
 1. lower atmosphere is evenly mixed: troposphere
 - a. nitrogen: 78%
 - b. oxygen: 21%
 - c. Argon nearly 1%
 - d. Water vapor variable up to 4%
 - 1) Dust and aerosols provide surfaces for condensation
 - 2) Condensation of water vapor into clouds releases latent heat
 - e. Carbon dioxide 0.037%--increasingly released from geologic storage
 2. stratosphere above troposphere, marked by different temperature
 - a. major gases about the same as troposphere

- b. contains most of ozone in atmosphere 20-30 km from surface
 - 1) ozone form of oxygen with three atoms, not two
 - 2) different properties:
 - a) absorbs UV light from Sun: our natural sunscreen: highly 'greenhouse' property
 - b) damaging to organism tissues
 - 3. outside the stratosphere, gases divided by molecular weight
- IV. Structure of atmosphere
 - A. Pressure diminishes with altitude
 - 1. Half of molecules are below 5.6 km
 - 2. Few are further than 16 km (10%)
 - B. Temperature changes more complicated
 - 1. troposphere—turning over part
 - a. heated from below, by reradiation from Earth
 - b. lower density with altitude accounts for some of cooler temperature—fewer molecules to run into one another
 - c. changes on average 6.5° C per 1000 m—environmental lapse rate
 - d. abrupt break where temperature stops decreasing--zone above troposphere is tropopause
 - 2. stratosphere—layer above troposphere: about 10 km
 - a. temperature stable, then increases gradually
 - b. most ozone is here: absorbs UV—causes the heating
 - 3. outer temperature zones
 - a. mesosphere decreases again—no ozone
 - b. thermosphere increase—direct radiation from Sun
- V. Atmospheric heating
 - A. Heat transfer
 - 1. radiation
 - 2. conduction
 - a. through matter by molecular activity
 - b. air is poor conductor of heat energy
 - 3. convection—heat rises
 - a. transfer of heat by circulation within fluid
 - b. caused by expansion of heated fluid reduces density
 - c. lower atmosphere heated by Earth's surface
 - B. radiation
 - 1. electromagnetic spectrum
 - a. gamma rays (short) to radio waves (long)
 - b. visible light small short range
 - c. ultraviolet shorter: sunburn rays
 - d. infrared longer: heat waves
 - 2. all bodies emit radiant energy
 - a. hotter bodies shorter
 - b. cooler bodies longer
 - 3. emit at their temperature wavelength

- a. Sun emits short—UV and visible
 - b. Earth emits longer—heat: energy absorbed from incoming solar radiation
 - C. Gases of atmosphere selectively absorb or transmit radiation, depending on wavelength
 - 1. mostly transparent to visible
 - 2. absorbs much UV and shorter wavelengths in outer atmosphere: oxygen and ozone
 - 3. infrared absorbed by carbon dioxide and water vapor
 - a. infrared from emission of energy absorbed from Sun
 - b. heats atmosphere from below by radiation and conduction
 - D. about ½ of incoming solar radiation absorbed by Earth's surface
 - 1. Albedo: reflection from clouds, snow, water—30%
 - 2. Scattering splits rays into numerous weaker rays: some to space
 - 3. Absorption by atmospheric gases
 - a. Greenhouse effect
 - 1) glass does not transmit heat, but also rely on lack of mixing with outside air to keep them warm
 - 2) greenhouse effect of atmosphere makes life on Earth possible
 - 3) Water vapor and carbon dioxide instrumental greenhouse effect gases
 - b. Oxygen and ozone absorb most of incoming UV wavelengths
 - E. Mostly between Tropical circles, where direct rays fall
 - 1. seasonal variations in Sun energy from constant axial tilt toward Polaris
 - a. solstice when direct rays most poleward—June 21 and December 21
 - b. equinox when direct rays at equator—March 21 and September 21
 - c. day length varies between extremes at solstices, equal on equinox
 - 2. additional causes of local temperature other than day length and Sun angle
- VI. Temperature
- A. Measure
 - 1. daily minimum and maximum are recorded
 - 2. these used to calculate all other statistics
 - a. Daily range: difference between high and low
 - b. Daily average: mean of the two values
 - c. Monthly average: mean of the daily averages for the month
 - d. Annual Mean: mean of the monthly averages
 - e. Annual range: difference between highest monthly average and lowest monthly averages
 - 3. maps of temperature plot 'isotherms'
 - a. lines of equal temperature

- 1) close—large differences across area
 - 2) far—small difference across area
 - b. useful for comparing areas, or comparing different seasons
 - 1) seasonal shifts of high temperature
 - 2) land water contrasts notable
 - 3) ocean currents can affect temperature
- B. Controls of temperature
 - 1. latitude
 - a. Sun angle
 - 1) High: more energy per unit area
 - 2) Low: less energy per unit area
 - b. Day length—greater variance at high latitude
 - c. Amount of atmosphere to penetrate—greater amount at high latitude
 - 2. land and water
 - a. water properties significant to heating differences
 - 1) water has high specific heat—can absorb much more energy with minimal change in temperature compared to land
 - 2) water is transparent—more than just surface is heated
 - 3) water is mobile—mixes cooler water with warmer
 - 4) water evaporates—
 - a) heat of vaporization 540 calories per gram!
 - b) Heat energy goes with the evaporated molecules
 - c) Released upon condensation of the water
 - b. Water moderates the temperature of areas
 - 1) Lower high temperatures
 - 2) Higher low temperatures
 - 3. Altitude
 - a. Lapse rate 6.5° C less per 1000 meters up
 - b. Some reradiation minimizes this lapse rate
 - 4. geographic position
 - a. leeward or windward coast
 - 1) get weather from sea—milder
 - 2) get weather from land—harsher
 - b. barriers—mountains block moisture
 - 5. cloud cover
 - a. clouds reflect incoming radiation—albedo in day
 - b. clouds absorb heat—greenhouse effect at night

VII. World temperature distribution

- A. All of the above affect distribution—latitude most pronounced
 - 1. sun angle
 - 2. day length
- B. Land-water contrast contributes to annual range
 - 1. coldest and hottest far from water
 - 2. most mild in -coastal tropical areas
- C. oceanic currents transport heat from equatorial area toward poles

- D. Maps are 'adjusted' for elevation (by lapse rate calculations)
 - 1. useful for comparing different times of year
 - 2. useful for noting anomalies, leading to investigations

VIII. Atmospheric moisture

- A. Water changes phase from solid to liquid to vapor with transfer of heat between environment and water molecules
 - 1. this transfer of heat in evaporation/condensation moves heat as 'latent heat'—hidden as energy is moved by water vapor
 - a. latent heat of vaporization cools environment as it evaporates
 - b. latent heat of condensation warms environment when droplets form
 - 2. sublimation and deposition are the change between solid and vapor
 - a. ice cubes can evaporate by sublimation
 - b. frost can form from vapor in the air
- B. saturation with water—humidity
 - 1. amount of water the air can hold
 - a. ways to measure
 - 1) mixing ratio—grams of water per kilogram of air
 - 2) relative humidity—percent of water compared to how much it could hold: more commonly reported, easier to measure
 - b. changing relative humidity
 - 1) adding moisture
 - a) above water bodies, especially hot springs
 - b) plants, animals also contribute
 - 2) warm air can hold more moisture
 - a) constant mixing ratio, increasing temperature will result in decrease of relative humidity
 - b) cooling temperature, constant mixing ratio results in increase of relative humidity
 - i. cools until air is saturated—100% relative humidity
 - ii. condensation begins
 - a. latent heat is released to environment
 - b. temperature stops decreasing
 - iii. this is the DEW POINT TEMPERATURE
 - 3) relative humidity changes throughout the day due to temperature change, with little change in mixing ratio
 - 2. measuring relative humidity
 - a. comparing air temperature to depression created by evaporation the air can accomplish with its moisture content
 - b. noting temperature at which condensation of water in a container that is gradually cooled